



2 The influence of climate state variables on Atlantic Tropical Cyclone 3 occurrence rates

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5 Received 1 January 2007; revised 20 April 2007; accepted 12 June 2007; published XX Month 2007.

6 [1] We analyzed annual North Atlantic tropical cyclone (TC) counts from 1871-2004,
7 considering three climate state variables—the El Niño/Southern Oscillation (ENSO),
8 peak (August-October or ‘ASO’) Sea Surface Temperatures (SST) over the main
9 development region (‘MDR’: 6-18°N, 20-60°W), and the North Atlantic Oscillation
10 (NAO)—thought to influence variations in annual TC counts on interannual and longer
11 timescales. The unconditional distribution of TC counts is observed to be inconsistent
12 with the null hypothesis of a fixed rate random (Poisson) process. However, using two
13 different methods, we find that conditioning TC counts on just two climate state variables,
14 ENSO and MDR SST, can account for much or all of the apparent non-random variations
15 over time in TC counts. Based on statistical models of annual Atlantic TC counts
16 developed in this study and current forecasts of climate state variables, we predicted
17 $m = 15 \pm 4$ total named storms for the 2007 season.

18 **Citation:** Sabbatelli, T. A., and M. E. Mann (2007), The influence of climate state variables on Atlantic Tropical Cyclone occurrence
19 rates, *J. Geophys. Res.*, 112, XXXXXX, doi:10.1029/2007JD008385.

21 1. Introduction

22 [2] A number of past studies have examined climatic
23 influences on variations at interannual and longer timescales
24 in the occurrence and the intensity of North Atlantic
25 Tropical Cyclones (TCs) [e.g., Gray, 1984]. The primary
26 factor considered in past studies is the El Niño/Southern
27 Oscillation (ENSO) [e.g., Bove et al., 1998; Landsea et al.,
28 1999; Elsner et al., 2000; Elsner, 2003; Elsner et al., 2006;
29 Elsner and Jagger, 2006], though the influence of the North
30 Atlantic Oscillation (‘NAO’) has also been examined in
31 some studies [Elsner et al., 2000; Elsner, 2003; Elsner et
32 al., 2006; Elsner and Jagger, 2006]. Both phenomena are
33 believed to influence TC production, development, or
34 prevailing trajectories through their influence on storm
35 tracks or vertical wind shear in the tropical North Atlantic.
36 The ENSO phenomenon tends to enhance (diminish) TC
37 counts during storm seasons coinciding with an incipient La
38 Nina (El Niño) event, while the NAO tends to enhance
39 (diminish) TC counts during storm seasons coinciding with
40 an incipient negative (positive) phase winter. Influences are
41 historically found only during the storm season preceding
42 the anomaly in the index; there is no detectable impact on
43 the following year’s storm season.

44 [3] Sea Surface Temperatures (SST) over the main
45 development region (‘MDR’: 6-18N, 20-60W) for North
46 Atlantic TCs during the season (August-October or ‘ASO’)
47 of Peak TC production [Emanuel, 2005a; Webster et al.,

2005, 2006; Mann and Emanuel, 2006; Srivier and Huber, 48
2006; Elsner, 2006] have also been argued to be an 49
important influence on long-term North Atlantic TC behav- 50
ior. MDR SSTs are considered a proxy for potential TC 51
intensity [Emanuel, 2005a], with annual TC counts 52
enhanced (diminished) in seasons associated with positive 53
(negative) MDR SST anomalies. Related studies have 54
argued for a significant influence of the so-called ‘Atlantic 55
Multidecadal Oscillation’ (‘AMO’) on North Atlantic TC 56
numbers [e.g., Goldenberg et al., 2001]. However, as the 57
procedures used to define the ‘AMO’ signal in terms of 58
North Atlantic SSTs in such studies has been challenged 59
in recent work [Trenberth and Shea, 2006; Mann and 60
Emanuel, 2006], we have chosen in our analyses here to 61
employ MDR ASO SSTs themselves [as in e.g., Emanuel, 62
2005; Mann and Emanuel, 2006; Elsner, 2006], rather than 63
an index such as the ‘AMO’ derived through statistical 64
processing of the North Atlantic SST field. 65

[4] Previous studies have investigated long-term trends in 66
TC statistics [e.g., Solow and Moore, 2000] or have used 67
regression models employing climatic indices [Gray, 1984; 68
Elsner et al., 2000, 2006; Elsner and Jagger, 2006] and trend 69
parameters [Elsner, 2003] to predict interannual variations in 70
TC activity. In no previous studies we are aware of, however, 71
have investigators examined whether conditioning on cli- 72
matic factors can account for the entirety of non-random 73
structure in the statistical distribution of historical North 74
Atlantic annual TC counts. In this study we perform such 75
an examination, employing two distinct and complementary 76
methods to test the hypothesis that annual TC counts follow a 77
state-dependent Poisson process against the null hypothesis 78
of a constant rate Poisson random process. 79

[5] Any statistical approach to analyzing TC counts must 80
respect the Poisson distributional nature of the underlying 81

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82 process (that is, that TC counts are characterized by a point
83 process with a low occurrence rate). Our first approach
84 employs Poisson regression [see e.g., *Elsner et al.*, 2000,
85 2001; *Elsner*, 2003; *Elsner and Jagger*, 2006], a variant on
86 linear regression which is appropriate for modeling a
87 conditional Poisson process in which the expected occur-
88 rence rate co-varies with some set of state variables (e.g.,
89 indices of ENSO, the NAO, and MDR SST). The second
90 approach categorizes the data with respect to the climate
91 state variables using a binary classification scheme, testing
92 both for the statistical significance of differences in occur-
93 rence rates between the resulting data subgroups, and
94 examining the resulting subgroup distributions for consis-
95 tency with a Poisson random process. The two methods are
96 complementary in that the latter method avoids the restric-
97 tive linearity assumptions implicit in regression, while the
98 former method accounts for continuous variations in
99 expected TC occurrence rates as a function of the underly-
100 ing state variables (e.g., distinguishing between the impacts
101 of strong vs. weak El Nino events).

102 2. Data

103 [6] Our analysis employed four data sets including (1)
104 historical annual North Atlantic TC counts, (2) the Decem-
105 ber-February (DJF) Niño3.4 SST ENSO index, (3) the
106 December-March (DJFM) NAO index, and (4) Aug-Oct
107 (ASO) seasonal SST means over the main development
108 region ('MDR') of 6°–18°N, 20°–60°W. Our analysis was
109 confined to the 135 year interval 1870–2004 over which all
110 three primary data sets of interest were available. The more
111 recent seasons of 2005 and 2006 for which preliminary data
112 are available, are subsequently interpreted in the context of
113 these analyses, while forecasts for the 2007 season are made
114 based on projected values of the climate indices. Data are
115 available at the supplementary website: [http://www.
116 meteo.psu.edu/~mann/TC_JGR07](http://www.meteo.psu.edu/~mann/TC_JGR07).

117 [7] Historical estimates of the annual TC counts are
118 available back to 1850 [*Jarvinen et al.*, 1984]. The reliabil-
119 ity of these data, particularly prior to the late 20th century in
120 which satellite and aircraft reconnaissance are available, has
121 been vigorously debated in recent studies [e.g., *Landsea*,
122 2005; *Emanuel*, 2005b]. *Emanuel* [2005b] nonetheless
123 makes a credible argument for why long-term TC count
124 data should be reliable, even if TC intensity estimates are
125 not. As *Emanuel* [2005b] notes, prior to aircraft reconnais-
126 sance, ships crossing the Atlantic would not have been
127 warned off from a developing or approaching storm, and
128 were likely to encounter either the storm or evidence of its
129 existence. Combined with other impacts on islands or
130 coastal localities, the existence of an Atlantic tropical
131 cyclone was therefore likely to have been known, even
132 prior to aircraft reconnaissance.

133 [8] Various alternative indices of the El Nino/Southern
134 Oscillation (ENSO) are available. We employed the boreal
135 winter (DJF) Niño3.4 index (SST averaged over the region
136 5°S–5°N, 120°–170°W) favored by many investigators [e.g.,
137 *Trenberth*, 1997]. Use of alternative (e.g., Niño3) ENSO
138 indices yielded similar conclusions. The Niño3.4 index was
139 taken from the *Kaplan et al.* [1998] data set and updated
140 with subsequent values available through NCEP. The boreal
141 winter (DJFM) NAO index was taken from *Jones et al.*

[1997], updated with more recent values from the Univer- 142
sity of East Anglia/CRU. For simplicity, the 'year' was 143
defined to apply to the preceding storm season for both 144
indices (e.g., the 1997/1998 El Nino and winter 1997/1998 145
NAO value were assigned the year 1997). 146

[9] The MDR SST index was taken from the HadISST2 147
observational SST data set [*Rayner et al.*, 2003] and 148
updated with more recent values from the UK Met Office. 149
The data were averaged over the season most relevant to 150
tropical cyclone formation (August–September–October, or 151
'ASO'). Estimated uncertainties in the observational SST 152
data are relatively small back to 1870 for both the Niño3.4 153
and North Atlantic regions of interest in this study [see, e.g., 154
Kaplan et al., 1998]. 155

3. Methods

[10] As in previous studies [e.g., *Elsner et al.*, 2000], we 157
assumed that annual TC counts n can be modeled as a 158
(Poisson) point process, viz. 159

$$P_i(n) = (1/n!) \mu^n \exp(-\mu) \quad (1)$$

where the mean occurrence rate μ , is the sole free parameter of 161
the distribution, and in the unconditional case has a Maximum 162
Likelihood value equal to the mean annual count. While the 163
appropriate null hypothesis holds the rate parameter μ to be 164
constant over time, it is of interest to investigate the 165
alternative hypothesis that μ may vary with respect to some 166
set of governing factors or 'state variables' [e.g., time—*Solow*
167 *and Moore*, 2000; *Elsner*, 2003 and/or climate state indices—
168 e.g., *Elsner*, 2003; *Elsner and Jagger*, 2006]. 169

[11] For the purposes of our study, μ was conditioned on 170
the three climate state variables discussed above (ENSO as 171
measured by the DJF Niño3.4 index, NAO as measured by, 172
the DJFM NAO index, and MDR SST as measured by the 173
MDR ASO SST index). Two distinct statistical approaches 174
were taken, as described below. We note that here is room 175
for further development of the methods presented below. 176
For example, one could extend the approaches used in the 177
present study to account explicitly for the increased uncer- 178
tainty in TC counts back in time, and in particular the 179
impact of unreported events [e.g., as in *Solow and Moore*,
180 2000; *Elsner and Jagger*, 2006]. 181

3.1. Binary Classification Approach

[12] In this approach, each year is classified as belonging 182
to one of two possible binary states (positive or negative) 183
with respect to each state variable, depending on sign of the 184
anomaly in that variable (relative to the 1870–2004 mean). 185
An alternative tertiary classification procedure was tested in 186
which a third neutral category was introduced (defined by 187
absolute anomalies within one standard deviation). The 188
choice of binary vs. tertiary classification schemes repre- 189
sents a tradeoff between the level of discrimination (two vs. 190
three states) and resulting sample sizes. While similar 191
results were obtained using the tertiary categorizations 192
scheme, we preferred the binary classification scheme due 193
to the larger sizes of the data sub-samples. For similar 194
reasons, only the two most significant (see section 4 for 195
further discussion) of the three state variables, MDR SST 196
and Niño3.4 were used. 197
198
199

[13] Using the binary classification scheme, we categorized years with respect to each of the two factors separately, and further, into three distinct sub-groupings, defined as (1) ‘favorable’: years in which both factors are favorable to TC production (positive MDR SST and negative Niño3.4 anomalies), (2) ‘unfavorable’: years in which both factors are unfavorable to TC production (negative MDR SST and positive Niño3.4 anomalies), and (3) ‘neutral’: years in which the two factors tend to offset in terms of their favorability to TC formation, i.e., anomalies in MDR SST and Niño3.4 that are of the same sign.

[14] We used a χ^2 test to evaluate the goodness-of-fit of a Poisson distribution for both the unconditional (i.e., all 135 years grouped together) and conditional (i.e., ‘favorable’, ‘neutral’, and ‘unfavorable’) data categorizations. We assumed χ^2 to have $\nu = B - 2$ degrees of freedom, where B is the number of occupied bins, and 2 degrees of freedom are subtracted based on constraints provided from the data (normalization of the distribution, and estimation of the rate parameter μ). The bin bandwidth was chosen using the objective criterion cited by *Wilks* [2005],

$$h \approx cIQR/N^{1/3} \quad (2)$$

where N is the sample size, IQR is the inter-fourth quartile range of the data, and $c = 2$ is taken for relatively skew distributions such as the Poisson. h was rounded to the nearest integer value.

[15] The t statistic was then used to evaluate the statistical significance of the differences in TC rate parameter estimates μ_i between any two data sub-samples. The t statistic reduces to

$$t = (\mu_1 - \mu_2) / (\mu_1/\phi_1 + \mu_2/\phi_2)^{1/2} \quad (3)$$

using the expression for the sample variance of a Poisson distribution, $\sigma^2 = \mu$, where ϕ_1 and ϕ_2 denote the degrees of freedom in the respective sub-samples, and the degrees of freedom in the t statistic is $\min(\phi_1, \phi_2) - 1$. When only Niño3.4—which is serially uncorrelated—is used as a conditioning variable, ϕ_1 and ϕ_2 reduce to simply N_1 and N_2 , the nominal sizes of the respective sub-samples. However, significant serial correlation in the MDR SST series (the lag one autocorrelation coefficient $\rho = 0.55$ yields a decorrelation timescale $\tau = 1.67$ years) decreases the effective number of independent climate states sampled when conditioning on MDR SST as, e.g., two neighboring years are not statistically independent with respect to the enhanced likelihood of elevated TC counts. Reduced degrees of freedom (ϕ) were therefore taken into account in estimating the statistical significance of t scores when conditioning fully or partly on the MDR SST series. In such cases, only events spaced more than two decorrelation timescales (i.e., 3 years) apart were considered to constitute statistically independent samples.

[16] Finally, we used a cross-validation procedure to evaluate the predictive skill in the binary conditional Poisson model approach. One could [see, e.g., *Elsner and Jagger*, 2006] leave each year out one at a time, forming conditional TC rate parameter estimates based on the remaining years and evaluating the skill of the resulting

classifications applied to each choice of missing year. However, when serial correlation is present in the state variables, which as discussed above is the case here, the results of such a cross-validation procedure are likely to give a too liberal an estimate of skill. We therefore employed an alternative split calibration/validation procedure. Conditional TC rate parameter estimates were obtained using the first half (i.e., years 1870-1937) of the data, and subsequently used to categorize the subsequent TC count data based on the climate state variable anomalies (measured relative to the calibration period baseline) over the latter half (i.e., years 1943-2004). This procedure was then repeated with the role of the first and last half of the data sets reversed. The average of the mean squared error (MSE) between the predicted and observed TC count data obtained for both sub-intervals was used as an estimate of cross-validated MSE, which was compared to the MSE obtained over the full (1870-2004) model development interval.

3.2. Poisson Regression

[17] Poisson regression is a variant on linear regression appropriate for data such as TC counts for which the null hypothesis of a Poisson distribution is appropriate [see *Elsner et al.*, 2000, 2001; *Elsner*, 2003; *Elsner and Jagger*, 2006 for further discussion]. Given a count series Y with unconditional mean rate μ believed to follow a state-dependent Poisson distribution, Poisson regression estimates a generalized linear model for the conditional expected rate of occurrence $\lambda = E(Y)$ as a function of a set of state variables X_1, X_2, \dots, X_M , of the form,

$$\log \lambda = \beta_0 + \beta_1 X_1 + \beta_2 X_2 + \dots + \beta_M X_M \quad (4)$$

or alternatively,

$$\lambda = \exp[\beta_0 + \beta_1 X_1 + \beta_2 X_2 + \dots + \beta_M X_M] \quad (5)$$

where the residuals are assumed to be Poisson distributed.

[18] Unlike ordinary linear regression, a closed-form analytical solution to equation (5) is not possible. However, it is straightforward to numerically estimate maximum likelihood values for the regression parameters β_i , and thus an estimate for the conditional expected occurrence rates λ_i . The residual series $\varepsilon_i = Y_i - \lambda_i + \mu$ can be analyzed for consistency with a Poisson distribution based on a χ^2 test, as described in section 3.1 above.

[19] Poisson regression was performed for various combinations of climate state variables as discussed in more detail in section 4. Cross-validation was performed using the split calibration/validation procedure discussed in section 3.1 wherein the regressions were performed alternatively using the first and last half of the full data set, with TC counts predicted and compared with observed counts over the remaining independent half of the data set. Quality of regression fit was measured by both the coefficient of determination R^2 and mean square error (MSE).

4. Results

[20] Certain relationships between annual TC counts and the Niño3.4 and MDR SST time series are evident by

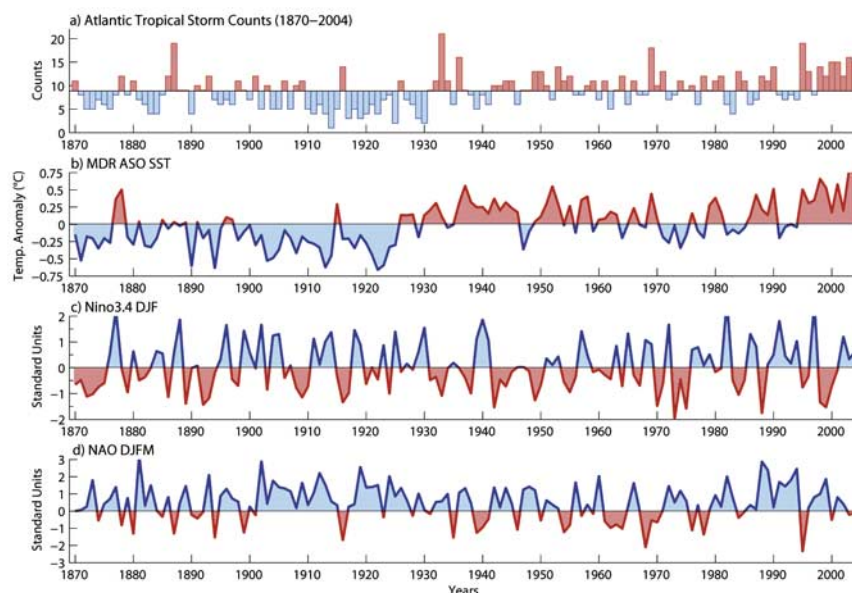


Figure 1. Time Series (1870-2004) of (a) annual Atlantic TC counts, (b) MDR ASO SST time series, (c) Niño3.4 DJF SST index, and (d) NAO DJFM SLP index. Red (blue) indicates positive (negative) anomalies in TC counts and Hurricane-favorable (unfavorable) conditions in the three indices (MDR SST, Niño3.4 and NAO). Note that year convention applies to the ‘D’ in DJF and DJFM for both ‘c’ and ‘d’.

inspection alone (Figure 1). The clear increase in TC counts subsequent to the 1920s, and the positive trend over roughly the past decade, closely coincide with corresponding tendencies for positive MDR SST anomalies. Anomalously low TC counts in certain years (e.g., 1982 and 1997) correspond to prominent El Niño years, and the low TC counts of the early 1990s correspond to general tendency for El Niño-like conditions. The NAO has a weaker, but nonetheless statistically significant impact on TC counts, with a tendency for elevation of counts during negative NAO years. The Pearson correlation coefficients between the TC counts and the three predictors ($r = 0.48$ for MDR SST, $r = -0.32$ for Niño3.4, and $r = -0.25$) are statistically significant at the $p < 0.0001$, $p = 0.0001$, and $p = 0.003$ levels respectively for a two-sided hypothesis test, taking into account the serial correlation in each series. The extent to which these state variables can account for the non-random structure in long-term TC counts is investigated below using each of the two methods discussed in section 3. Figure 2.

4.1. Binary Classification Approach

[21] We first note that the unconditional distribution of TC counts is highly inconsistent with the null hypothesis of a random Poisson process. Based on a χ^2 test (Table 1) we reject at the $p < 0.05$ level the null hypothesis of a Poisson process for the entire TC count record 1870-2004. By inspection (Figure 3, panel a), it is clear that there is bimodality in the distribution which cannot be captured by the model of a constant mean Poisson process.

[22] Conditioning on ENSO influences (i.e., on Niño3.4) alone does not ameliorate this problem, as the conditional distributions for negative Niño3.4 values (i.e., ‘La Niña’-like behavior) is still observed (Table 1) to be inconsistent

($p < 0.05$) with a Poisson distribution. Conditioning on MDR SST provides significant improvement, though the values ($p = 0.79$ and $p = 0.25$ for +MDR SST and -MDR SST respectively) average only just above the median ($p = 0.5$) level between acceptance and rejection of the null hypothesis. However, when TC counts are simultaneously conditioned on both Niño3.4 and MDR SST, we find that the null hypothesis can likely not be rejected. The resulting three separate distributions (‘favorable’, ‘neutral’, and ‘unfavorable’, as defined in section 3.1) are generally well captured by a Poisson distribution (Figure 3, panels b-d). While in one of the three cases (‘favorable’) the p value ($p = 0.27$) indicates a moderate 27% chance of falsely rejecting the null hypothesis, the χ^2 tests yield an average value $p = 0.70$ for the three cases, well above the median expected level for false rejection of the null hypothesis. The results of the analysis are therefore consistent with the hypothesis that the annual TC counts are produced by a state-dependent Poisson process, with the occurrence rate being dictated by two state variables (Niño3.4 and MDR SST).

[23] Having established the viability of a state-dependent Poisson random model for the observed TC count data, we assessed the statistical significance of differences in the estimated conditional occurrence rates μ . There is a clear dependence of μ both on each of the two state variables separately and on the sub-categorization into the three ‘favorable’, ‘neutral’, and ‘unfavorable’ cases (Table 2). The highest average annual TC count is found for the ‘favorable’ state ($\mu \approx 11$), while the lowest ($\mu \approx 6$) is found for the ‘unfavorable’ state, with all other sub-groupings yielding intermediate values of μ . While differences in occurrence rate (Table 3) are highly significant conditioning on either one of the two state variables (Niño3.4 or MDR SST) alone, the most significant difference (i.e., lowest p

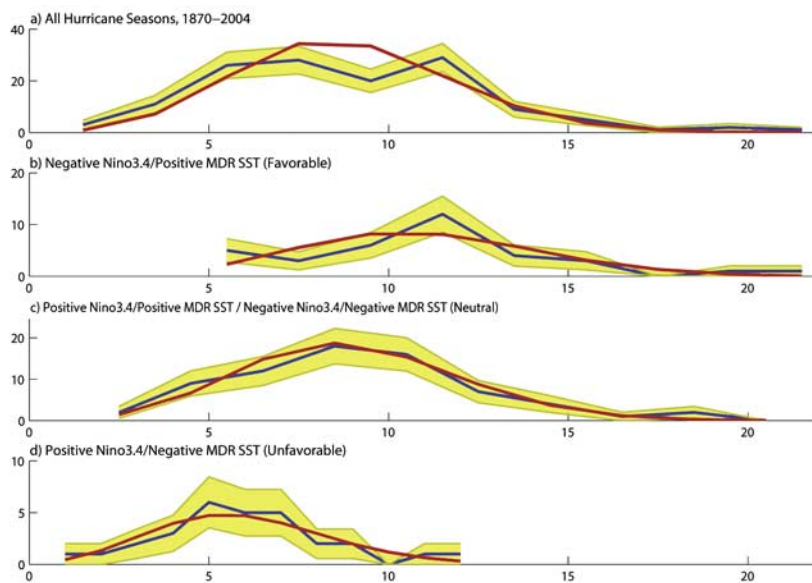


Figure 2. Histograms of TC counts n vs. bin centers (blue) with associated one standard deviation uncertainties ($\pm\sqrt{n}$, yellow shading) and best fit Poisson distributions (red). Results are shown for unconditional case (all data—panel *a*) and the ‘favorable’, ‘neutral’, and ‘unfavorable’ sub-groupings discussed in the text (panels *b-d*). Bin bandwidths were determined as discussed in text.

379 value) is observed conditioning on both state variables (i.e.,
 380 the ‘unfavorable’ vs. ‘favorable’ categories). Partitioning
 381 into the ‘favorable’, ‘neutral’, and ‘unfavorable’ categories
 382 yields both individual distributions that as noted earlier are
 383 on average consistent with Poisson, and mean TC occur-
 384 rence rates that differ significantly between any two cate-
 385 gories (Table 3). The MSE (Table 4) using the conditional
 386 means from the binary classification approach (MSE =
 387 10.80 for the full 1870-2004 model development interval,
 388 and MSE = 11.79 in cross-validation) represents a signifi-
 389 cant improvement over climatology (MSE = 13.75) or
 390 persistence (MSE = 19.89). The cross-validation results,
 391 however, suggest that the binary classification approach
 392 gives moderately less predictive skill than the Poisson
 393 regression approach, as discussed in more detail below.

394 4.2. Poisson Regression

395 [24] We performed univariate Poisson regression alterna-
 396 tively using (i) MDR SST and (ii) Niño3.4 as state variables,
 397 (iii) bivariate regression using both MDR SST and Niño3.4
 398 as state variables, and (iv) multivariate regression using all
 399 three climate state variables MDR SST, Niño3.4, and NAO
 400 (Figure 3a). Cross-validated resolved variance R^2 and MSE
 401 scores were similar to the scores obtained from the full
 402 model development interval 1870-2004, and far superior to
 403 either climatology or persistence, indicating significant skill
 404 in each of the regression models. Interestingly, the predic-
 405 tive skill systematically increases while the consistency of
 406 residuals (see Figure 3b) with a Poisson distribution
 407 decreases as additional state variables are added to the
 408 regression—i.e., first MDR only, then MDR and Niño3.4,
 409 and finally MDR, Niño3.4 and NAO (Table 4). Improved

skill thus appears to come at a cost of increased bias in the 410
 conditional TC rate estimates. 411

[25] Each of the Poisson regression models are seen to 412
 improve significantly (as measured by both full 1870-2004 413
 model development interval and cross-validation MSE 414
 scores) over climatology (Table 4). Moreover, both bivariate 415
 and three variable Poisson regression models yield signifi- 416
 cant improvements (as measured by MSE scores) over the 417
 binary classification approach with MDR SST and Niño3 418
 outlined in section 4.1. This further suggests a tendency for 419
 a tradeoff between resolved variance (as determined from 420
 regression and validation R^2 and MSE scores) and bias (as 421
 determined from the distribution of residuals) in modeling 422
 TC counts. While the binary classification approach yielded 423
 the greatest consistency with a pure state-dependent Poisson 424
 process (as conditional distributions were consistent with 425
 Poisson at a mean level $p = 0.70$), it also produced the least 426
 resolved variance in modeling annual TC counts by condi- 427
 tioning on two or more climate state variables. 428

Table 1. Results of Reduced χ^2 Tests Described in Text^a

Scenario (1870-2004)	χ^2/ν	ν	p	t1.1
All Years	2.09	9	0.027	t1.3
+MDR SST	0.59	8	0.79	t1.4
−MDR SST	1.32	3	0.25	t1.5
+Niño3.4	1.02	8	0.42	t1.6
−Niño3.4	2.29	7	0.025	t1.7
+MDR/−Niño (‘Favorable’)	1.27	6	0.27	t1.8
−MDR/+Niño (‘Unfavorable’)	0.28	9	0.98	t1.9
+MDR/+Niño or −MDR/−Niño (‘Neutral’)	0.49	7	0.85	t1.10

^aIndicated are reduced χ^2 value (χ^2/ν), degrees of freedom ν and the p value for rejection of the null hypothesis of a poisson distribution.

t1.11

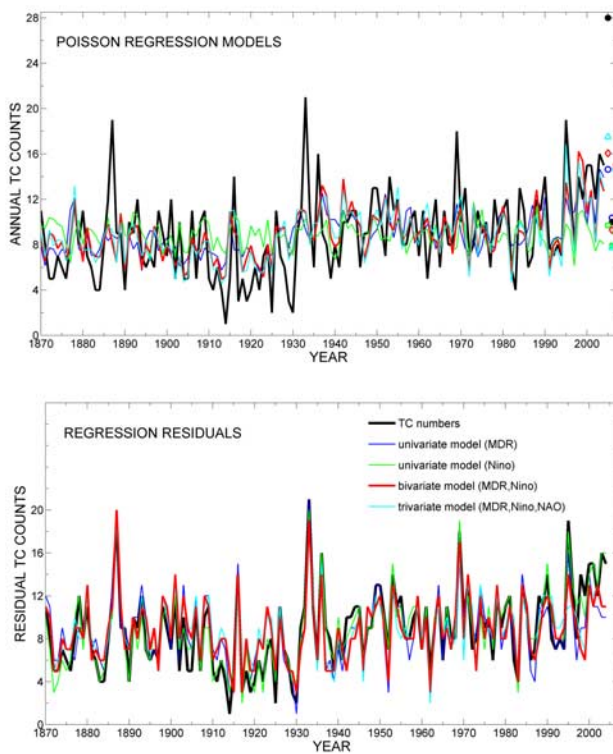


Figure 3. Poisson regression models of annual Atlantic TC counts using the MDR ASO SST, Niño3.4, and NAO series as predictors. Shown are (a) the statistical model fits over 1870-2004 based on the two univariate, bivariate and three-variable Poisson regressions (colored curves) along with the observed TC counts for 1870-2004 (black curve), observed TC counts for 2005 and 2006 (filled black circles), predicted TC counts for 2005 and 2006 (unfilled colored symbols) and 2007 (filled colored symbols). (b) Poisson regression residuals as defined in text (colored curves) along with the observed TC counts for 1870-2004 (black curve).

429 **4.3. Predictions**

430 [26] The binary classification approach to modeling TC
 431 numbers yields a simple forecasting scheme for seasonal TC
 432 counts. Depending on the forecast values for the two state
 433 variables (MDR ASO SST and DJF Niño3.4 anomalies) at
 434 the start of the tropical cyclone season (June 1st), the
 435 predicted TC total would be $\mu = 6 \pm 3$ (i.e., between 3

Table 3. Results of t Tests for Differences of Occurrence Rates μ Among the Different Sub-Groupings Discussed in Text^a

Scenario (1870-2004)	t	Φ	P
+MDR SST vs. -MDR SST	3.59	27	0.0006
+Nino3.4 vs. -Nino3.4	3.70	57	0.0002
Favorable vs. Unfavorable	5.41	19	<0.0001
Favorable vs. Neutral	2.15	19	0.02
Neutral vs. Unfavorable	4.02	19	0.0004

^aIndicated are the effective degrees of freedom in the t statistic $\Phi = \min(\phi_1, \phi_2) - 1$, and the one-tailed p value for rejection of the null hypothesis of equal means.

and 9) for ‘unfavorable’ anomaly combinations, $\mu = 9 \pm 3$ (between 6 and 12) for ‘neutral’ anomaly combinations, and $\mu = 11 \pm 3$ (between 8 and 14) for ‘favorable’ anomaly combinations. It is instructive to interpret the two most recent (2005 and 2006) Atlantic tropical storm seasons in this context. The TC count for the 2006 season ($n = 10$) was consistent with the predicted count ($m = 9 \pm 3$) given the observed ‘neutral’ conditions (positive MDR SST anomaly and positive 2006/2007 DJF Niño3.4 anomaly—see Table 5). The 2005 TC count ($n = 28$) is considerably more difficult to explain, even given the ‘favorable’ (positive 2005 MDR SST and negative 2005/2006 DJF Niño3.4) observed conditions, for which the predicted count is $m = 11 \pm 3$. Given a mean expected rate $\mu = 11$, the probability of equaling or exceeding a TC count of $n = 28$ is $\approx 0.01\%$, i.e., implausible.

[27] The Poisson regression models all successfully predict the 2006 TC count within estimated uncertainties, but like the binary classification approach, all significantly under-predict the historic 2005 TC total of $n = 28$ storms (Table 5, and also Figure 3a). However, the most skillful of the Poisson regression models as judged by cross-validation results (i.e., Table 3)—the three state variable model—comes closest to the observed total with a predicted TC count of $m = 18 \pm 4$. The high predicted total in this case is a result of simultaneously favorable conditions in all three state variables (anomalously warm MDR ASO SSTs, La Nina conditions in the tropical Pacific, and a substantially negative phase NAO). Given a conditional expected mean

Table 4. Assessments of Predictive Skill for Competing Statistical Models Considered in This Study^a

Model/Predictors	R^2_{full}	MSE_{full}	R^2_{valid}	MSE_{valid}	p_{resid}
Climatology	0.00	13.75			
Persistence	0.07	19.89			
Binary Cond: MDR, Nino		10.80		11.79	
Poisson Reg: MDR	0.24	10.81	0.16	10.47	0.83
Poisson Reg: Nino	0.10	12.51	0.12	12.31	0.08
Poisson Reg: MDR, Nino	0.33	9.37	0.26	9.95	0.35
Poisson Reg: MDR, Nino, NAO	0.38	8.70	0.32	9.02	0.00

^aMean square error (MSE) over the full model development period (1870-2004) is indicated for each case. The MSE for simple (i) climatological mean and (ii) persistence predictions is provided for comparison. In the case of Poisson regression models, the coefficient of determination (R^2) is also provided. Validation MSE and R^2 scores are based on the split calibration/validation procedures described in the text.

Table 2. Estimates of Occurrence Rate μ for the Various TC Data Sub-Groupings Discussed in Text^a

Scenario (1870-2004)	μ	N	ϕ
All Years	8.85	135	
+MDR SST	10.33	64	28
-MDR SST	7.52	71	31
+Nino3.4	7.78	58	
-Nino3.4	9.66	77	
+Nino/+MDR (‘Favorable’)	10.94	35	20
-MDR/+Nino (‘Unfavorable’)	5.97	29	20
+MDR/+Nino or -MDR/-Nino (‘Neutral’)	9	71	33

^aProvided are the sample sizes N and, where appropriate, the effective sample size ϕ accounting for temporal autocorrelation in state variables.

t5.1 **Table 5.** Climate State Variable Values and Associated Annual TC Count Predictions m and Associated one Standard Error Uncertainties $\pm\sqrt{m}$ for 2005-2007^a

t5.2	Year	Model	MDR	Nino3.4	NAO	Predicted (n)	Observed (m)
t5.3	2005	Binary conditioning	+	-	x	11 ± 3	28
t5.4		Poisson regression	x	-0.65	x	10 ± 3	
t5.5			28.87C	x	x	15 ± 4	
t5.6			28.87C	-0.65	x	16 ± 4	
t5.7			28.87C	-0.65	-0.82	18 ± 4	
t5.8	2006	Binary conditioning	+	+	x	9 ± 3	10
t5.9		Poisson regression	x	0.72	x	8 ± 3	
t5.10			28.35C	x	x	10 ± 3	
t5.11			28.35C	0.72	x	9 ± 3	
t5.12			28.35C	0.72	2.43	8 ± 3	
t5.13	2007	Binary conditioning	+	-	x	11 ± 3	To be determined
t5.14		Poisson regression	x	-0.2	x	10 ± 3	
t5.15			27.9C ^b	x	x	15 ± 4	
t5.16			27.9C ^b	-0.2 ^b	x	15 ± 4	
t5.17			27.9C ^b	-0.2 ^b	0.47 ^b	15 ± 4	
t5.18	^a 2007 climate variables are forecast based on the procedure described in the text.						
t5.19	^b Predicted value.						

480 rate $\mu = 18$, the probability of observing or exceeding $n = 28$
 481 storms is approximately 2%. In other words, for every
 482 50 years with conditions similar to those observed for
 483 2005, a TC count as high or higher than that observed
 484 might be expected given the three variable Poisson regres-
 485 sion model. In this case, the 2005 TC total is still observed
 486 to be improbable, but not entirely implausible. It is of course
 487 possible that the true distribution of TC occurrence is
 488 heavy-tailed, in which case the probability of very large
 489 counts might be substantially greater than estimated
 490 under the assumption of conditional Poisson statistics.
 491 One could conceivably also argue that biases in the earlier
 492 data [e.g., Landsea, 2005] leads to an underestimation of the
 493 frequency of very large annual counts such as observed in
 494 2005. However, our finding in section a that long-term TC
 495 data are essentially consistent with random Poisson statistics
 496 after controlling for dependence on two climate state
 497 variables, would seem to argue against the proposition that
 498 systematic biases compromise the reliability of the earlier
 499 data [Landsea, 2005].

500 [28] Finally, we use the statistical models developed
 501 above to forecast Atlantic TC counts for the 2007 tropical
 502 storm season. At the time this manuscript was finalized,
 503 weak La Nina conditions (Nino3.4 = -0.2) were predicted
 504 by NCEP for winter 2007/2008. MDR SST anomalies were
 505 currently similar to those observed for the 2005 season, so
 506 we infer by persistence ASO MDR SST anomalies equal to
 507 those for the 2005 season. As there is no basis for forecast-
 508 ing the winter 2007/2008 NAO value, we assume climatolo-
 509 gical mean DJFM conditions (NAO index = 0.47). Given
 510 these assumed values, the binary classification approach
 511 yields the ‘favorable’ forecast $m = 11 \pm 3$, while each of
 512 the Poisson regression models (with the exception of the
 513 Niño3.4-only regression which yields a forecast $m = 11 \pm 3$)
 514 predict a total of $m = 15 \pm 4$ storms for the 2007 tropical
 515 storm season.

517 5. Conclusions

518 [29] Two different methods, a binary classification
 519 scheme and Poisson regression, are used to condition

520 expected annual TC counts on climate state variables.
 521 Modeling annual Atlantic TC counts as a state-dependent
 522 Poisson process using the binary classification approach, we
 523 find that two climatic factors, ENSO and tropical North
 524 Atlantic MDR SST, are adequate to explain the apparent
 525 non-random variability in historical variations in Atlantic
 526 TC numbers. Modeling TC counts instead using Poisson
 527 regression, we find that the most skillful statistical model
 528 employs all three state variables considered in the study,
 529 ENSO, tropical North Atlantic MDR SST, and the NAO, as
 530 predictors. This three variable statistical model also comes
 531 closest to predicting the historic 2005 TC count of 18,
 532 ascribing unlike the other statistical models developed in
 533 this study, a non-trivial probability for that event given the
 534 climate state of 2005. However, analysis of residuals also
 535 indicates some evidence of bias, implying the need for
 536 cautious use of the model. Three of the four Poisson
 537 regression models developed in the study predict 15 ± 4
 538 storms for the 2007 Atlantic tropical storm season.

[30] **Acknowledgments.** We thank two anonymous reviewers for
 their helpful comments. We also acknowledge generous financial support
 (for T.S.) provided by the Pennsylvania State University Schreyer Honor
 College.

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